Biomass offsets little or none of permafrost carbon release from soils, streams, and wildfire: an expert assessment.

Benjamin W. Abbott1,2\*, Jeremy B. Jones2, Edward A.G. Schuur3, F. Stuart Chapin III2, William B. Bowden4, M. Syndonia Bret-Harte2, Howard E. Epstein5, Michael D. Flannigan6, Tamara K. Harms2, Teresa N. Hollingsworth7, Michelle C. Mack3, A. David McGuire8, Susan M. Natali9, Adrian V. Rocha10, Suzanne E. Tank11, Merritt R. Turetsky12, Jorien E. Vonk13, Kimberly P. Wickland14, George R. Aiken14, Heather D. Alexander15, Rainer M.W. Amon16, Brian W Benscoter17, Yves Bergeron18, Kevin Bishop19, Olivier Blarquez20, Ben Bond-Lamberty21, Amy L. Breen22, Ishi Buffam23, Yihua Cai24, Christopher Carcaillet25, Sean K. Carey26, Jing M. Chen27, Han Y. H. Chen28, Torben R. Christensen29, Lee W. Cooper30, J. Hans C. Cornelissen31, William J. de Groot32, Thomas H. DeLuca33, Ellen Dorrepaal34, Ned Fetcher35, Jacques C. Finlay36, Bruce C. Forbes37, Nancy H.F. French38, Sylvie Gauthier39, Martin P. Girardin39, Scott J. Goetz9, Johann G. Goldammer40, Laura Gough41, Paul Grogan42, Laodong Guo43, Philip E. Higuera44, Larry Hinzman45, Feng Sheng Hu46, Gustaf Hugelius47, Elchin E. Jafarov48, Randi Jandt49, Jill F. Johnstone50, Jan Karlsson51, Eric S. Kasischke52, Gerhard Kattner53, Ryan Kelly54, Frida Keuper55, George W. Kling56, Pirkko Kortelainen57, Jari Kouki58, Peter Kuhry59, Hjalmar Laudon60, Isabelle Laurion61, Robie W. Macdonald62, Paul J. Mann63, Pertti J. Martikainen64, James W. McClelland65, Ulf Molau66, Steven F. Oberbauer67, David Olefeldt68, David Paré39, Marc-André Parisien69, Serge Payette70, Changhui Peng71, Oleg S. Pokrovsky72, Edward B. Rastetter73, Peter A. Raymond74, Martha K. Raynolds75, Guillermo Rein76, James F. Reynolds77, Martin Robards78, Brendan M. Rogers9, Christina Schädel3, Kevin Schaefer79, Inger K. Schmidt80, Anatoly Shvidenko81, Jasper Sky82, Robert G.M. Spencer83, Gregory Starr84, Robert G. Striegl14, Roman Teisserenc85, Lars J. Tranvik86, Tarmo Virtanen87, Jeffrey M. Welker88, Sergei Zimov89

1.Université de Rennes 1, OSUR, CNRS, UMR 6553 ECOBIO, 2. Institute of Arctic Biology and Department of Biology & Wildlife, University of Alaska Fairbanks, 3. Center for Ecosystem Science and Society, Northern Arizona University, 4. The Rubenstein School of Environment and Natural Resources, University of Vermont, 5. Department of Environmental Sciences, University of Virginia, 6. Department of Renewable Resources, University of Alberta, 7. USDA Forest Service, PNW Research Station; University of Alaska Fairbanks, 8. U.S. Geological Survey, Alaska Cooperative Fish and Wildlife Research Unit; University of Alaska Fairbanks, 9. Woods Hole Research Center, 10. Department of Biological Sciences and the Environmental Change Initiative, University of Notre Dame, 11. Department of Biological Sciences, University of Alberta, 12. Department of Integrative Biology, University of Guelph, 13. Department of Earth Sciences, Utrecht University, 14. US Geological Survey, National Research Program, Boulder, CO, 15. University of Texas-Brownsville, 16. Texas A&M University at Galveston, 17. Florida Atlantic University, 18. Forest Research Institute, Université du Québec en Abitibi-Témiscamingue, 19. Uppsala University, Department of Earth Sciences; Department of Aquatic Sciences and Assessment, Swedish University of Agricultural Sciences, 20. Département de Géographie, Université de Montréal, 21. Pacific Northwest National Lab, 22. Scenarios Network for Alaska & Arctic Planning, International Arctic Research Center, University of Alaska Fairbanks, 23. University of Cincinnati, 24. State Key Laboratory of Marine Environmental Science, Xiamen University, 25. Ecole Pratique des Hautes Etudes, UMR5023 CNRS Lyon 1, 26. McMaster University, 27. University of Toronto, 28. Lakehead University, 29. Lund University, Arctic Research Centre; Aarhus University, 30. University of Maryland Center for Environmental Science, 31. Systems Ecology, VU University, 32. Canadian Forest Service, Natural Resources Canada, 33. School of Environmental and Forest Sciences, University of Washington, 34. Climate Impacts Research Centre, Umeå University, 35. Institute for Environmental Science and Sustainability, Wilkes University, 36. Department of Ecology, Evolution and Behavior, University of Minnesota, 37. Arctic Centre, University of Lapland, 38. Michigan Tech Research Institute, Michigan Technological University, 39. Natural Resources Canada, Canadian Forest Service, Laurentian Forestry Centre, 40. Global Fire Monitoring Center, Max Planck Institute for Chemistry, 41. Department of Biological Sciences, Towson University, 42. Department of Biology, Queen's University, 43. University of Wisconsin-Milwaukee, School of Freshwater Sciences, 44. Department of Ecosystem and Conservation Sciences, University of Montana, 45. University of Alaska Fairbanks, 46. Department of Plant Biology and Department of Geology, University of Illinois, 47. Department of Physical Geography, Stockholm University, 48. Arctic and Alpine Research, University of Colorado Boulder, 49. Alaska Fire Science Consortium, University of Alaska Fairbanks, 50. Biology Department, University of Saskatchewan, 51. Climate Impacts Research Centre Abisko, Department of Ecology and Environmental Science, Umeå University, 52. Department of Geographical Sciences, University of Maryland, 53. Alfred Wegener Institute Helmholtz Center for Polar and Marine Research, 54. Neptune and Company Inc. 55. Climate Impacts Research Centre Abisko, Umeå University; INRA, AgroImpact UPR1158, 56. University of Michigan, 57. Finnish Environment Institute, 58. School of Forest Sciences, University of Eastern Finland, 59. Department of Physical Geography, Stockholm University, 60. Department of Forest Ecology and Management, Swedish University of Agricultural Sciences, Umeå, 61. Centre Eau Terre Environnement, Institut national de la recherche scientifique, 62. Institute of Ocean Science, 63. Department of Geography, Northumbria University, 64. Department of Environmental and Biological Sciences, University of Eastern Finland, 65. University of Texas at Austin, Marine Science Institute, 66. University of Gothenburg, Department of Biological and Environmental Sciences, 67. Florida International University, Department of Biological Sciences, 68. University of Alberta, Department of Revewable Resources, 69. Natural Resources Canada, Canadian Forest Service, Northern Forestry Centre, 70. Centre d’études Nordiques, Université Laval, 71. Center of CEF/ESCER, University of Quebec at Montreal; State Key Laboratory of Soil Erosion and Dryland Farming on the Loess Plateau, College of Forestry, Northwest A&F University, 72. Georesources and Environment, CNRS, Toulouse; BIO-GEO-CLIM Laboratory, Tomsk State University, 73. The Ecosystems Center, Marine Biological Laboratory, Woods Hole, 74. Yale School of Forestry and Environmental Studies, 75. Institute of Arctic Biology, University of Alaska Fairbanks, 76. Department of Mechanical Engineering, Imperial College London, 77. School of Life Sciences, Lanzhou University; Nicholas School of the Environment, Duke University, 78. Wildlife Conservation Society, Arctic Beringia Program, 79. National Snow and Ice Data Center, Cooperative Institute for Research in Environmental Sciences, University of Colorado Boulder, 80. Department of Geosciences and Natural Resource Management, University of Copenhagen, 81. International Institute for Applied Systems Analysis, Laxenburg, Austria; Sukachev Institute of Forest, 82. Affiliate, Cambridge Centre for Climate Change Research, 83. Florida State University, Department of Earth, Ocean & Atmospheric Science, 84. Department of Biological Sciences, University of Alabama, 85. ECOLAB, Université de Toulouse, CNRS, INPT, UPS, 86. Limnology, Department of Ecology and Genetics, Uppsala University, 87. Department of Environmental Sciences, University of Helsinki, 88. University of Alaska Anchorage, 89. Northeast Science Station of the Russian Academy of Sciences

\* benabbo@gmail.com, Campus de Beaulieu, bât. 15, 263 av du Général Leclerc, 35042 Rennes CEDEX – France, Tel +33750911654, Fax +33223236077,

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**Abstract**

As the permafrost region warms, its large organic carbon pool will be increasingly vulnerable to decomposition, combustion, and hydrologic export. Models predict that some portion of this release will be offset by increased production of Arctic and boreal biomass; however, the lack of robust estimates of net carbon balance increases the risk of further overshooting international emissions targets. Precise empirical or model-based assessments of the critical factors driving carbon balance are unlikely in the near future, so to address this gap, we present estimates from 98 permafrost-region experts of the response of biomass, wildfire, and hydrologic carbon flux to climate change. Results suggest that contrary to model projections, total permafrost-region biomass could decrease due to water stress and disturbance, factors that are not adequately incorporated in current models. Assessments indicate that end-of-the-century organic carbon release from Arctic rivers and collapsing coastlines could increase by 75% while carbon loss via burning could increase four-fold. Experts identified water balance, shifts in vegetation community, and permafrost degradation as the key sources of uncertainty in predicting future system response. In combination with previous findings, results suggest the permafrost region will become a carbon source to the atmosphere by 2100 regardless of warming scenario but that 65 to 85% of permafrost carbon release can still be avoided if human emissions are actively reduced.

**Introduction**

**Permafrost zone carbon balance**

The United Nations has set a target of limiting warming to 2°C above pre-industrial temperatures to mitigate risk of the most damaging consequences of climate change ([UNEP, 2013](#_ENREF_75)). Maintaining global climate within this target depends on understanding ecosystem feedbacks to climate change so that adequate limits on human emissions can be set. As high latitudes warm, more of the large permafrost carbon pool will be exposed to decomposition, combustion, and hydrologic export ([Harden *et al.*, 2012](#_ENREF_25); [Schuur *et al.*, 2015](#_ENREF_69)). Up to 220 Petagrams (Pg) carbon could be released from permafrost-region soil by 2100, and 500 Pg by 2300 ([Schuur *et al.*, 2013](#_ENREF_68); [MacDougall *et al.*, 2012](#_ENREF_48)), representing 10 to 30% of greenhouse gas emissions required to push the global climate system beyond the 2°C target ([Schaefer *et al.*, 2014](#_ENREF_65)). Models project that some permafrost carbon release will be offset by increases in Arctic and boreal primary productivity due to extended growing season, CO2 fertilization, and nutrient release from decomposing soil organic matter. However, many processes and dynamics known to influence biomass accumulation, such as ecosystem disturbance and nutrient limitation, are incompletely represented or absent in current models ([Qian *et al.*, 2010](#_ENREF_60); [Koven *et al.*, 2011](#_ENREF_40); [Schaefer *et al.*, 2011](#_ENREF_66); [Koven *et al.*, 2015b](#_ENREF_41)). Likewise, only a few models projecting future permafrost carbon release consider wildfire emissions, and none include hydrologic carbon flux ([MacDougall *et al.*, 2012](#_ENREF_48); [Koven *et al.*, 2011](#_ENREF_40); [Qian *et al.*, 2010](#_ENREF_60); [Schaefer *et al.*, 2014](#_ENREF_65); [Schaefer *et al.*, 2011](#_ENREF_66)),though past hydrologic flux has been simulated ([McGuire *et al.*, 2010](#_ENREF_51); [Kicklighter *et al.*, 2013](#_ENREF_37); [Laudon *et al.*, 2012](#_ENREF_43)). Despite clear policy implications of this climate feedback, considerable uncertainty of both carbon inputs and outputs limits our ability to model carbon balance of the permafrost region. To bring to bear the best available quantitative and qualitative scientific information ([Joly *et al.*, 2010](#_ENREF_34)) on this climate feedback, we present results from expert assessment surveys indicating that there is little consensus on the magnitude and even sign of change in high-latitude biomass, whereas most researchers expect fire emissions and hydrologic organic carbon flux to substantially increase by the end of the century.

**Expert assessment**

When data are sparse but management decisions are pressing, expert judgements have long been used to constrain possible system response and risk of dangerous or undesired outcomes ([Zickfeld *et al.*, 2010](#_ENREF_83); [Morgan, 2014](#_ENREF_52)). There are multiple methods for collecting and combining expert opinion including formal expert elicitation interviews, interactive software, and surveys ([Aspinall, 2010](#_ENREF_5); [Morgan, 2014](#_ENREF_52); [Javeline *et al.*, 2013](#_ENREF_33)). While expert assessment cannot definitively answer questions of future system response, it complements modeling and empirical approaches by allowing the synthesis of formal and informal system information and by identifying research priorities (Fig. 1;([Morgan, 2014](#_ENREF_52); [Sutherland *et al.*, 2013](#_ENREF_72)). The approach is similar to the concept of ensemble models where multiple estimates built on different assumptions and data provide a more robust estimate and measure of variance. Because the experimental unit is an individual researcher, each data point represents an integration of quantitative knowledge from modeling, field, and laboratory studies as well as qualitative information based on professional opinion and personal experience with the system. Expert assessment has been used in risk assessment and forecasting of natural disasters, human impacts on ecosystems, and tipping points in the climate system ([Aspinall, 2010](#_ENREF_5); [Halpern *et al.*, 2008](#_ENREF_24); [Lenton *et al.*, 2008](#_ENREF_46)). In a data-limited environment such as the permafrost region, expert assessment allows formal consideration of a range of factors known to affect carbon balance but insufficiently quantified for inclusion in models. For permafrost carbon balance, these factors include nutrient dynamics, non-linear shifts in vegetation community, human disturbance, land-water interactions, and the relationship of permafrost degradation with water balance.

Because precise empirical or model-based assessments of the critical factors driving permafrost-region carbon balance are unlikely in the near future ([Harden *et al.*, 2012](#_ENREF_25)), we collected estimates of the components of net ecosystem carbon balance from 98 permafrost-region experts (Table 1). We had two major goals: 1. Assess current understanding of the timing and magnitude of non-soil biomass accumulation, hydrologic organic carbon flux, and wildfire carbon emissions, and 2. Identify major sources of uncertainty in high-latitude carbon balance to inform future research.

**Methods**

**Survey development and design**

In the fall of 2013 we administered three expert assessments to address knowledge gaps concerning the response of permafrost-region biomass, wildfire, and hydrologic carbon flux to climate change. Development of assessment methodology began in early 2009 as a part of the Dangerous Climate Change Assessment Project administered by the University of Oxford. We iteratively revised questions, response format, and background information based on four rounds of input from participants, including at the Vulnerability of Permafrost Carbon Research Coordination Network meeting in Seattle 2011 ([Schuur *et al.*, 2013](#_ENREF_68)). To help survey participants consider all of the evidence available from field and modeling studies, we distributed a system summary document for each questionnaire including regional and pan-Arctic estimates of current carbon pools and fluxes, a brief treatment of historical trends, and a summary of model projections where available (Table 2; Supplementary information Questionnaires and System summaries).

Participants were selected based on contribution to peer-reviewed literature or referrals from other experts and had experience in all major boreal and Arctic regions (Table 1). We identified potential participants by querying Thomas Reuters Web of Science (webofknowledge.com) with applicable search terms (e.g. Arctic, boreal, biomass, dissolved organic carbon, fire, permafrost). To reach researchers with applicable expertise who were underrepresented in the literature, we supplemented the list with personal referrals from lead experts and all participants. In total 256 experts were invited to participate. We distributed the surveys and system summaries via email with a two-week deadline. After sending out three reminders and accepting responses for three months after initial invitation, we received 115 responses from 98 experts (38% response rate), with 15 experts participating in more than one survey (Supplementary information List of experts). Experts who provided estimates and input to this paper are coauthors.

Experts provided quantitative estimates of change in biomass, hydrologic flux, or wildfire for three time points (2040, 2100, and 2300), and four regional warming scenarios based on representative concentration pathway (RCP) scenarios from the IPCC Fifth Assessment Report ([Moss *et al.*, 2010](#_ENREF_53)). Warming scenarios ranged from cessation of human emissions before 2100 (RCP2.6) to sustained human emissions (RCP8.5) and corresponded to permafrost-region mean annual warming of 2 to 7.5°C by 2100. All surveys were driven by the same scenarios of high-latitude warming generated from RCP2.6, 4.5, 6.0, and 8.5 with the National Center for Atmospheric Research's Community Climate System Model 4 ([Lawrence *et al.*, 2012](#_ENREF_45)). For the purposes of this survey, warming was assumed to stabilize at 2100 levels for all scenarios so that responses through 2300 accounted for lags in ecosystem responses to climate drivers. While climate scenarios were defined by temperature, we asked experts to consider all accompanying direct climate effects (e.g. temperature, precipitation, and atmospheric CO2) and indirect effects (e.g. vegetation shifts, permafrost degradation, invasive species, and disturbance). Experts were encouraged to consider all available formal and informal information when generating their estimates including published and unpublished modeled and empirical data as well as professional judgment. Participants listed the major sources of uncertainty in their estimates, self-rated their confidence and expertise for each question, described rationale for their estimates, and provided background information (Tables 1 and S1).

The biomass survey consisted of a single question asking for cumulative change in tundra and boreal non-soil biomass including above and belowground living biomass, standing deadwood, and litter. The wildfire survey asked for estimates of change in wildfire extent and CO2 emissions for the boreal and tundra regions to assess changes in both fire extent and severity. The hydrologic flux survey asked for estimates of dissolved and particulate organic carbon (DOC and POC, respectively) delivery to freshwater ecosystems in the pan-Arctic watershed and delivery to the Arctic Ocean and surrounding seas via riverine flux and coastal erosion, allowing the calculation of losses during transport due to burial or mineralization. Dissolved inorganic carbon fluxes were not included in this survey.

The original questionnaires in 2009 asked for participants to estimate subjective 95% confidence intervals of the whole system response (e.g. total change in high-latitude biomass). Based on expert input during subsequent testing we disaggregated the system into different components to encourage detailed consideration of possibly competing dynamics ([Morgan, 2014](#_ENREF_52)) (e.g. asking for separate estimates of boreal forest and Arctic tundra response). This resulted in a large response table for each question (72-102 quantitative estimates), which we found caused respondent fatigue and decreased the number of experts willing to participate. As a compromise, we asked respondents to provide a single best estimate and indicate confidence with a five-point scale (Table S1). While analysis of best estimates can return narrower uncertainty ranges than subjective probability distributions ([Morgan, 2014](#_ENREF_52)), we believe this tradeoff resulted in broader expert participation, better representing diversity of opinion across disciplines and compensating for possible underestimation of variability and uncertainty.

**Analysis and calculations**

We calculated basic summary statistics, using median values to estimate center and interquartile ranges (IQR) to estimate spread. To calculate the portion of permafrost carbon release offset by biomass accumulation, we combined estimates from this study with reanalyzed data from Schuur *et al.* (2013). The low IQR for carbon release offset by biomass growth was calculated by dividing the low IQR of uptake by the upper IQR of carbon release and conversely for the high IQR. All analyses were performed in R 3.0.2. The complete dataset of quantitative estimates and comments from survey participants stripped of personal identifiers is available at [www.aoncadis.org/dataset/Permafrost\_carbon\_balance\_survey.html](http://www.aoncadis.org/dataset/Permafrost_carbon_balance_survey.html).

**Results**

**Carbon pools and fluxes**

Expert estimates revealed diverging views on the response of boreal biomass to warming, with over a third of estimates predicting a decrease or no change in boreal biomass across scenarios and time periods (Fig. 2). While median change in boreal biomass was similar across warming scenarios for each time step (3, 9, and 11% increases by 2040, 2100, and 2300, respectively; Figs. 2 and S1), variability was much higher for warmer scenarios. Consequently, all of the interquartile ranges of change in boreal biomass for RCP6.0 and RCP8.5 included zero. Experts projecting a decrease in boreal biomass attributed their estimates primarily to water-stress and disturbance such as fire and permafrost degradation. In contrast, there was general agreement that tundra biomass would respond positively to warming, with end-of-century increases of 6 to 30% projected for RCP2.6 and 10 to 90% for RCP8.5. Because of these contrasting responses to increased warming, tundra accounted for 40% of total biomass gain by 2300 for RCP8.5, though it currently constitutes less than 10% of total permafrost region biomass (based on median values in Fig. 2; Fig. 3a; Table 2). Estimates of boreal biomass were generally symmetrically distributed while tundra biomass estimates were right-skewed, and most datasets had 1 to 4 estimates beyond 1.5 times the interquartile range (Fig. S2). Self-rated confidence was higher for tundra than for boreal forest, but was below 3 (moderately confident) in both cases (Table S1), highlighting considerable uncertainty of individual estimates in addition to variability among respondents.

Experts projected major shifts in both fire and hydrologic carbon regimes, with up to a 75% increase of riverine organic carbon flux to the ocean and a four-fold increase in fire emissions by 2100 for RCP8.5 based on interquartile ranges (Fig. 2 and S1). Fire and hydrologic carbon release estimates peaked at 2100, followed by a 10 to 40% decrease through 2300. In contrast to biomass, the response of both fire-driven and hydrologic carbon flux varied strongly by warming scenario, with RCP8.5 resulting in 2 to 6 times more carbon release than RCP2.6. While the boreal forest dominated total wildfire emissions, the relative change in tundra fire emissions was 1.5 and 2-fold greater than the relative boreal response for 2100 and 2300, respectively (Fig. S1). Changes in fire emissions were attributed to changes in fire extent rather than severity, which varied less than 5% among scenarios and time periods. Though dissolved organic carbon (DOC) represented the majority of total hydrologic organic carbon release, experts projected higher relative increases for coastal particulate organic carbon (POC), with end-of-the-century increases of 6 to 50% for RCP2.6 and 13 to 190% for RCP8.5. There was a lack of consensus on the response of DOC delivery to the ocean, with 21% of estimates predicting a decrease or no change. Experts predicting a decrease attributed their estimates to increased mineralization, changes in hydrologic flowpath, and changes in DOC photo- and bio-lability ([Cory *et al.*, 2014](#_ENREF_13); [Abbott *et al.*, 2014](#_ENREF_3)). Responses indicated no change in the proportion of organic carbon mineralized or trapped in sediment before reaching the ocean, with 63-69% of DOC and 68-74% of POC lost in transport. Fire and hydrologic carbon flux estimates were strongly right-skewed with a few experts projecting extreme change well beyond 1.5 times the interquartile range for each timestep and warming scenario combination (Figs. S3 and S4). Average self-rated confidence was between 2 and 3 for all questions except tundra fire emissions which had average confidence of 2.0 and 1.7 (Table S1).

**Sources of uncertainty**

Along with quantitative estimates of carbon balance, experts identified sources of uncertainty currently limiting the prediction of system response to climate change (Table 3). Water balance, including precipitation, soil moisture, runoff, infiltration, and discharge, was the most frequently mentioned source of uncertainty for both biomass and hydrologic organic carbon flux, and the second most mentioned for wildfire. Many experts noted that water balance is as or more important than temperature in controlling future carbon balance, yet projections of water balance are less well constrained ([Zhang *et al.*, 2013](#_ENREF_82); [Bintanja and Selten, 2014](#_ENREF_10)). Almost three-quarters of wildfire experts identified the future distribution of vegetation as the primary source of uncertainty in projecting wildfire, noting strong differences in flammability between different boreal and tundra species. Permafrost degradation was identified as an important source of uncertainty for biomass, hydrologic flux, and wildfire, due to both disturbance from ground collapse (thermokarst) and interactions with water-table dynamics and surface soil moisture as deeper thaw affects soil drainage.

**Discussion**

**Carbon balance**

Arctic tundra and boreal forest have accumulated a vast pool of organic carbon, twice as large as the atmospheric carbon pool and three times as large as the carbon contained by all living things ([Hugelius *et al.*, 2014](#_ENREF_30); [Schuur *et al.*, 2015](#_ENREF_69)). Over the past several decades, the permafrost region has removed an average of 500 Tg carbon yr-1 from the atmosphere ([McGuire *et al.*, 2009](#_ENREF_50); [Pan *et al.*, 2011](#_ENREF_56); [Hayes *et al.*, 2011](#_ENREF_26)). Combining our estimates of biomass uptake with a recent projection of permafrost soil carbon release ([Schuur *et al.*, 2013](#_ENREF_68)) suggests that the permafrost region will become a carbon source to the atmosphere by 2100 for all warming scenarios (Fig. 3b). Experts predicted that boreal and Arctic biomass could respond more quickly to warming than soil carbon release, offsetting -33 to 200% of mid-century emissions from permafrost-region soil (Fig. 3b). However, because estimates of change in biomass are similar across warming scenarios while permafrost carbon release is strongly temperature-sensitive, the emissions gap widens for warmer scenarios, resulting in 5-times more net carbon release under RCP8.5 than RCP2.6. This suggests that 65 to 85% of permafrost carbon release could be avoided if human emissions are actively reduced—*i.e.* if emissions follow RCP2.6 instead of RCP8.5 (Fig. 4).

**Comparison with quantitative models**

Model projections of future boreal and Arctic biomass agree in sign but vary widely in magnitude, with increases of 9 to 61 Pg carbon projected by 2100 ([Qian *et al.*, 2010](#_ENREF_60); [Koven *et al.*, 2011](#_ENREF_40); [Schaefer *et al.*, 2011](#_ENREF_66); [Falloon *et al.*, 2012](#_ENREF_17)). While some of these models fall within the range estimated here of -20 to 28 Pg carbon by 2100, none include zero or negative change in biomass as predicted by over a third of participants in our expert assessment. Two potential reasons for this disagreement are an overestimation of the effect of CO2 fertilization or an underestimation of the role of disturbance in some models. Firstly, CO2 fertilization exerts a larger effect on carbon balance than all other climate effects in many models ([Balshi *et al.*, 2009](#_ENREF_7)), with up to 88 Pg carbon difference between model runs with and without CO2 fertilization effects for some models ([Koven *et al.*, 2011](#_ENREF_40)). However, there is little field evidence that CO2 fertilization results in long-term biomass accumulation in tundra and boreal ecosystems ([Hickler *et al.*, 2008](#_ENREF_28); [Peñuelas *et al.*, 2011](#_ENREF_58); [Gedalof and Berg, 2010](#_ENREF_20)). Additionally, many models with large CO2 effects do not include other limiting factors, such as nutrients and water, known to interact with CO2 fertilization ([Hyvonen *et al.*, 2007](#_ENREF_31); [Yarie and Van Cleve, 2010](#_ENREF_80); [Thornton *et al.*, 2007](#_ENREF_73); [Koven *et al.*, 2015a](#_ENREF_39); [Maaroufi *et al.*, 2015](#_ENREF_47)). Secondly, models that do not account for disturbance such as wildfire, permafrost collapse, insect damage, and human resource extraction likely overestimate the positive response of biomass to climate change ([Kurz *et al.*, 2008](#_ENREF_42); [Abbott and Jones, 2015](#_ENREF_1); [Hewitt *et al.*, 2015](#_ENREF_27)).

Considering the scenario of a complete biome shift is useful in evaluating both model projections of change and estimates from our expert assessment. If all boreal forest became temperate forest, living biomass would increase by 27%, resulting in the uptake of 16 Pg carbon based on average carbon densities from both ecosystems ([Pan *et al.*, 2011](#_ENREF_56)). However, 22 Pg carbon would be lost due to decreases in dead wood and litter, resulting in a net circumboreal loss of 6 Pg carbon. If all tundra became boreal forest, non-soil biomass would increase by 205% ([Epstein *et al.*, 2012](#_ENREF_16); [Raynolds *et al.*, 2012](#_ENREF_61); [Saugier *et al.*, 2001](#_ENREF_64)), taking up 17 Pg carbon. This scenario may not represent the upper limit of possible carbon uptake if other unforeseen shifts in C allocation take place; however, it highlights the relatively modest carbon gains probable on century timescales.

While model regional projections of boreal wildfire vary in sign and magnitude (Supplementary information System summaries), most circumboreal models agree that fire emissions will increase several-fold, with increases of 200 to 560% projected by the end of the century ([Kloster *et al.*, 2012](#_ENREF_38); [Flannigan *et al.*, 2009](#_ENREF_18)). Interquartile ranges from our study are somewhat lower (40 to 300%, median 170%), but participant confidence in these estimates was low, suggesting considerable uncertainty in the future response of boreal fire. The 60 to 480% increase in tundra fire projected by our study would represent an even larger ecological shift than experienced by the boreal forest, with implications for regional biomass, habitat, and carbon balance, though there are few models that project changes in tundra fire ([Rupp *et al.*, 2000](#_ENREF_63)) and none at a circumarctic scale ([Mack *et al.*, 2011](#_ENREF_49)).

The production of Arctic DOC and POC depends on abundance of carbon sources in terrestrial ecosystems (influenced by biomass, wildfire, temperature, and permafrost degradation) and the ability of hydrologic flow to transport that carbon (determined by factors such as precipitation, runoff, depth of flow through soil, and coastal erosion) ([Guo *et al.*, 2007](#_ENREF_23); [Kicklighter *et al.*, 2013](#_ENREF_37); [Abbott *et al.*, 2015](#_ENREF_2)). Due to these complexities and others, there are currently no quantitative projections of future DOC and POC flux from the circumarctic. However, estimates from our study suggest a substantial departure from historical rates of change. For RCP8.5, hydrologic organic carbon loading would increase 4-20 times faster in the 21st century than it did in the 20th ([Kicklighter *et al.*, 2013](#_ENREF_37)), representing a non-linear response to high-latitude warming. The lack of consensus on the response of DOC, the largest component of hydrologic organic carbon flux, highlights the importance of developing and testing conceptual frameworks to be incorporated into models ([Laudon *et al.*, 2012](#_ENREF_43)).

An alternative explanation for differences between expert estimates and modeled projections is the possibility of bias in the group of experts. Participants in our assessment tended to have more field than modeling experience (Table 1) and may have therefore been skeptical of simulated ecosystem responses that have not been observed in the field such as CO2 fertilization and rapid migration of treeline ([McGuire *et al.*, 2009](#_ENREF_50)). Because future dynamics cannot always reliably be predicted on the basis of past system behavior, this bias may or may not result in overly conservative estimates. Furthermore, because experts are likely to base projections on the study areas with which they are most familiar, regional differences could be a source of bias. Fundamental differences among regions in the response of DOC flux and fire-regime to warming have been observed ([Kicklighter *et al.*, 2013](#_ENREF_37); [de Groot *et al.*, 2013](#_ENREF_14)); Supplementary information System summaries). Asia, which represents more than half of the total permafrost region, was under-represented in all three surveys, particularly wildfire (Table 1). However, the regional bias in this study may not be greater than that of model projections, which depend on observational and experimental data that are not evenly distributed throughout the permafrost region.

**Reducing uncertainty surrounding the permafrost carbon feedback**

Experts identified water balance, vegetation distribution, and permafrost degradation as the most important sources of uncertainty in predicting the timing and magnitude of the permafrost carbon feedback (Table 3). These three processes are closely interconnected by several internal feedbacks ([Jorgenson *et al.*, 2013](#_ENREF_36); [Shur and Jorgenson, 2007](#_ENREF_70); [Anisimov and Reneva, 2006](#_ENREF_4); [Girardin *et al.*, 2015](#_ENREF_22)). For example, wildfire or drought can trigger a transition from coniferous to deciduous dominance, warming permafrost by up to 7°C due to loss of insulating moss and associated changes ([Shur and Jorgenson, 2007](#_ENREF_70); [Sturm *et al.*, 2001](#_ENREF_71); [Yarie and Van Cleve, 2010](#_ENREF_80)). The subsequent recovery trajectories of vegetation and permafrost, as well as the proportion of thawed carbon released CO2 or CH4, then depend largely on near-surface hydrologic conditions ([Myers-Smith *et al.*, 2008](#_ENREF_54); [O'Donnell *et al.*, 2011](#_ENREF_55); [Jorgenson *et al.*, 2010](#_ENREF_35); [Chapin *et al.*, 2010](#_ENREF_12); [Lawrence *et al.*, 2015](#_ENREF_44); [Payette *et al.*, 2004](#_ENREF_57)). These interdepencies mean that improving projections of the permafrost carbon feedback will require conceptualizing these parameters together. The question of water balance is additionally important in Arctic and boreal ecosystems where hydrologic carbon flux can be the determining factor causing net carbon uptake or release (Oquist et al ; Kling et al). The lack of model projections of hydrologic carbon fluxes is a major gap in our ability to estimate the permafrost carbon feedback.

The permafrost region has responded differently to various climatic perturbations in the past, representing another tool to constrain possible future response ([Zachos *et al.*, 2008](#_ENREF_81)). During the Paleocene-Eocene Thermal Maximum, high-latitude temperature warmed more than 10°C, causing almost complete loss of permafrost and the mineralization of most permafrost soil organic matter ([DeConto *et al.*, 2012](#_ENREF_15); [Bowen and Zachos, 2010](#_ENREF_11)). More recently, the 2-4°C warming at high-latitudes during the early Holocene caused active-layer deepening throughout the permafrost region but did not trigger complete permafrost loss or widespread carbon release ([French, 1999](#_ENREF_19); [Schirrmeister *et al.*, 2002](#_ENREF_67); [Jorgenson *et al.*, 2013](#_ENREF_36)). While there are many differences between the Paleozoic and Holocene warming events, one clear distinction is the degree of warming. There may have been a threshold between 4 and 10°C high-latitude warming due to positive feedbacks such as a coniferous—deciduous shift or abrupt change in hydrology. If a tipping point does exist between 4 and 10°C high-latitude warming, it would fall between scenarios RCP4.5 and RCP8.5, representing maximum atmospheric CO2 of 650 and 850 ppm, respectively ([Moss *et al.*, 2010](#_ENREF_53); [Lawrence *et al.*, 2012](#_ENREF_45)). RCP4.5 is still widely accepted as politically and technically attainable, though it assumes global CO2 emissions peak before 2050 and decrease by half by 2080 ([Moss *et al.*, 2010](#_ENREF_53)).

**Conclusions**

The permafrost climate feedback has been portrayed in popular media (and to a lesser extent in peer-reviewed literature) as an all-or-nothing scenario. Permafrost greenhouse gas release has been described as a tipping point, a runaway climate feedback, and, most dramatically, a time bomb ([Wieczorek *et al.*, 2011](#_ENREF_79); [Treat and Frolking, 2013](#_ENREF_74); [Whiteman *et al.*, 2013](#_ENREF_78)). On the other extreme, some have dismissed the importance of this feedback, asserting that increases in biomass will offset any carbon losses from soil, or that changes will occur too slowly to concern current governments ([Idso *et al.*, 2014](#_ENREF_32)). Our expert estimates suggest that, while Arctic and boreal biomass may offset much or all of mid-century permafrost carbon release, the permafrost region will become a carbon source to the atmosphere by 2100 regardless of warming scenaio. However, results indicate a 5-fold difference in emissions between the business as usual scenario (RCP8.5) and active reduction of human emissions (RCP2.6), suggesting that up to 85% of carbon release from the permafrost region can still be avoided, though the window of opportunity for keeping that carbon in the ground is rapidly closing. Models projecting a strong boreal carbon sink and models that do not consider hydrologic and fire emissions may substantially underestimate net carbon release from the permafrost region. If such projections are used as the basis for emissions negotiations, climate targets are likely to be overshot.

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**Figure 1.** Conceptual model of the role of expert assessment in generating and communicating scientific understanding. Modelling and field research generate quantitative and qualitative understanding of the system (in this case the permafrost zone). Expert assessment synthesizes current understanding including qualitative information not yet included in numerical models. These syntheses provide perspective to the scientific community and wholistic summaries of the state of the knowledge to the non-scientific community with the goal of improving management of the system.

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**Figure 2.** Estimates of change in non-soil biomass, wildfire emissions, and hydrologic C flux from the permafrost region for four warming scenarios at three time points. All values represent change from current pools or fluxes reported in Table 2. Biomass includes above and belowground living biomass, standing deadwood, and litter. Dissolved and particulate organic C (DOC and POC respectively) fluxes represent transfer of C from terrestrial to aquatic ecosystems. "Coast" represents POC released by coastal erosion. For relative change see Fig. S1. Representative concentration pathway (RCP) scenarios range from aggressive emissions reductions (RCP2.6) to sustained human emissions (RCP8.5). Box plots represent median, quartiles, and minimum and maximum within 1.5 times the interquartile range. Full distributions are presented in Figs. S2 to S4.

**Figure 3.** Total change in non-soil biomass (a) and percentage of permafrost region C release offset by change in non-soil biomass (b). Estimates of permafrost C release used in estimating percentage offset are recalculated from data presented in Schuur et al. ([Schuur *et al.*, 2013](#_ENREF_68)). See Fig. 2 for definition of RCP scenarios and symbology. Error bars represent propogated error between the interquartile ranges of carbon release from permafrost soil and carbon uptake by biomass (see Methods).

**Figure 4.** A comparison of soil C release (recalculated from Schuur et al. ([Schuur *et al.*, 2013](#_ENREF_68))) and non-soil biomass uptake in the permafrost region for two warming scenarios. Polygons represent median cumulative change and dotted lines represent the interquartile range. Biomass C uptake is overlayed on soil C release to show the proportion of C release potentially offset by biomass. Linear rates of change were assumed between the three dates where estimates were provided. See Fig. 2 for definition of RCP scenarios.

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| --- | --- | --- | --- | --- |
| **Table 1.** Composition and characteristics of participant group | | | | |
|  | Biomass | | Wildfire | Hydrologic flux |
| Number of respondents | | 46 | 34 | 35 |
| Average responses per question\*  **Primary region of study** | | 41 | 28 | 32 |
| Asia | | 10 | 3 | 8 |
| Europe | | 12 | 5 | 9 |
| North America | | 27 | 27 | 18 |
| Circumpolar | | 12 | 6 | 9 |
| **Primary biome of study** | |  |  |  |
| Arctic | | 31 | 13 | 27 |
| Boreal | | 27 | 29 | 18 |
| Both | | 14 | 9 | 12 |
| Average modeling/field self rating\*\* | | 3.6 | 3.7 | 4.1 |
| Combined years of experience | | 762 | 533 | 521 |
| Ratio male:female | | 2.6 | 2.8 | 4.9 |

Background information on survey participants. Experts could indicate multiple regions and biomes of study. \*Not all experts provided estimates for all questions. \*\*Experts rated themselves on a 1-5 scale where 1=exclusive modeler and 5=exclusive field researcher.

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| --- | --- | --- | --- | --- | --- | --- | --- | --- | --- | --- | --- | --- | --- | --- | --- | --- | --- | --- | --- | --- | --- |
| Table 2. Estimates of current permafrost region organic carbon pools and fluxes | | | | | | | | | | | | | | | | | | | | | |
| **Biomass** | | | | | | | | | | | | | | | | | | | | | |
|  |  | | | Aboveground biomass | | Belowground biomassa | | | | | | | | Dead woodb | | | Litter | Total non-soil biomass | | | |
| Boreal forest (Pg C) | | | 43.6c | | 16.1 | | | | | | | | 16 | | | 27b | 102.7 | | | |
| Arctic Tundra (Pg C) | | | 2.4d | | 4.0 | | | | | | | |  | | | 2e | 8.4 | | | |
|  | |  |  | | | |  | | | | | |  | | | |  |  | | | |
| **Wildfire** | | | | | | | | | | | | | | | | | | | | | |
|  |  | | | | Boreal forest (Eurasia) | | | | | Boreal forest (N. America) | | | | | Total Boreal forestf | | | | | | Total Tundra |
| Area burned (km2 yr-1) | | | | 62,100 | | | | | 22,500 | | | | | 84,600 | | | | | | 4,200g |
| CO2 emissions from fire (Tg C yr-1) | | | | 194 | | | | | 56 | | | | | 250 | | | | | | 8h |
|  | |  | | |  | | | | | |  | | | | |  | | |  | | |
| **Hydrologic organic carbon flux** | | | | | | | | | | | | | | | | | | | | | |
|  |  | | | | | | | DOC | | | | POC (Riverine) | | | | POC (coastal) | | | | Total OC | |
| Delivery to freshwater ecosystems (Tg yr-1) | | | | | | | | 100c | | | 20i | | | | na | | | | 120 | |
| Delivery to Arctic Ocean and surrounding seas (Tg yr-1) | | | | | | | | 36j | | | 6c | | | | 18ck | | | | 60 | |

a ([Saugier *et al.*, 2001](#_ENREF_64)), b ([Pan *et al.*, 2011](#_ENREF_56)), c ([McGuire *et al.*, 2009](#_ENREF_50)), d ([Epstein *et al.*, 2012](#_ENREF_16)), e ([Potter and Klooster, 1997](#_ENREF_59)), f ([Balshi *et al.*, 2007](#_ENREF_8); [Giglio *et al.*, 2010](#_ENREF_21); [Hayes *et al.*, 2011](#_ENREF_26); [van der Werf *et al.*, 2010](#_ENREF_76)), g ([Rocha *et al.*, 2012](#_ENREF_62)), h ([Mack *et al.*, 2011](#_ENREF_49)), i ([Aufdenkampe *et al.*, 2011](#_ENREF_6); [Battin *et al.*, 2009](#_ENREF_9)), j ([Holmes *et al.*, 2012](#_ENREF_29)), k ([Vonk *et al.*, 2012](#_ENREF_77)). Literature-based estimates of belowground biomass were calculated from aboveground or total biomass with ratios from Saugier et al. ([Saugier *et al.*, 2001](#_ENREF_64)). POC delivery to freshwater ecosystems was calculated from ocean POC delivery with downscaled global ratio of 0.75 for sedimentation. POC from coastal erosion is the sum of Vonk et al. ([Vonk *et al.*, 2012](#_ENREF_77)) and McGuire et al. ([McGuire *et al.*, 2009](#_ENREF_50)). Considerable uncertainty remains around many of these estimates.

|  |  |  |  |  |  |
| --- | --- | --- | --- | --- | --- |
| Table 3. Sources of uncertainty in system response to climate change | | | | | |
| **Biomass** | | **Wildfire** | | **Hydrologic OC flux** | |
| Source of uncertainty | % | Source of uncertainty | % | Source of uncertainty | % |
| Water balance | 56 | Vegetation shift | 73 | Water balance | 41 |
| Wildfire | 47 | Water balance | 58 | Hydrologic flowpath | 39 |
| Permafrost degradation | 40 | Human disturbance | 27 | Permafrost degradation | 24 |
| Human disturbance | 29 | Permafrost degradation | 18 | Photo and bio-lability | 24 |
| Insect damage | 27 | Seasonality | 15 | Vegetation shift | 20 |
| Vegetation shift | 24 | Regional differences | 12 | Fluvial erosion | 11 |
| Treeline dynamics | 16 |  |  |  |  |
| Nutrient availability | 13 |  |  |  |  |
| Non-insect herbivores | 11 |  |  |  |  |

Major factors contributing uncertainty to projections of future system response based on expert comments. Rank is based on percent of experts who listed each factor in their responses. All sources listed by 10% or more of each group are included here. Water balance includes comments mentioning precipitation, soil moisture, runoff, infiltration, or discharge. Permafrost degradation includes comments referring to permafrost collapse (thermokarst) and active layer deepening.

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Figure 1

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Figure 2

Macintosh HD:Users:Ben:Dropbox:PHD:Abbott dissertation:CH2 NECB:February 2015 revision:Figures:Fig2.pdfFigure 3

Macintosh HD:Users:Ben:Dropbox:PHD:Abbott dissertation:CH2 NECB:February 2015 revision:Figures:Febfig3.pdfFigure 4

Abbott B W and Jones J B 2015 Permafrost collapse alters soil carbon stocks, respiration, CH4, and N2O in upland tundra *Glob. Change Biol.* **21** 4570-87

Abbott B W, Jones J B, Godsey S E, Larouche J R and Bowden W B 2015 Patterns and persistence of hydrologic carbon and nutrient export from collapsing upland permafrost *Biogeosciences* **12** 3725-40

Abbott B W, Larouche J R, Jones J B, Bowden W B and Balser A W 2014 Elevated dissolved organic carbon biodegradability from thawing and collapsing permafrost *Journal of Geophysical Research: Biogeosciences* 2014JG002678

Anisimov O and Reneva S 2006 Permafrost and changing climate: the Russian perspective *AMBIO: A Journal of the Human Environment* **35** 169-75

Aspinall W 2010 A route to more tractable expert advice *Nature* **463** 294-5

Aufdenkampe A K, Mayorga E, Raymond P A, Melack J M, Doney S C, Alin S R, Aalto R E and Yoo K 2011 Riverine coupling of biogeochemical cycles between land, oceans, and atmosphere *Frontiers in Ecology and the Environment* **9** 53-60

Balshi M S, McGuire A D, Duffy P, Flannigan M, Kicklighter D W and Melillo J 2009 Vulnerability of carbon storage in North American boreal forests to wildfires during the 21st century *Glob. Change Biol.* **15** 1491-510

Balshi M S, McGuire A D, Zhuang Q, Melillo J, Kicklighter D W, Kasischke E, Wirth C, Flannigan M, Harden J, Clein J S, Burnside T J, McAllister J, Kurz W A, Apps M and Shvidenko A 2007 The role of historical fire disturbance in the carbon dynamics of the pan-boreal region: A process-based analysis *J Geophys Res-Biogeo* **112** -

Battin T J, Luyssaert S, Kaplan L A, Aufdenkampe A K, Richter A and Tranvik L J 2009 The boundless carbon cycle *Nature Geosci* **2** 598-600

Bintanja R and Selten F M 2014 Future increases in Arctic precipitation linked to local evaporation and sea-ice retreat *Nature* **509** 479-82

Bowen G J and Zachos J C 2010 Rapid carbon sequestration at the termination of the Palaeocene-Eocene Thermal Maximum *Nature Geoscience* **3** 866-9

Chapin F S, McGuire A D, Ruess R W, Hollingsworth T N, Mack M C, Johnstone J F, Kasischke E S, Euskirchen E S, Jones J B, Jorgenson M T, Kielland K, Kofinas G P, Turetsky M R, Yarie J, Lloyd A H and Taylor D L 2010 Resilience of Alaska's boreal forest to climatic change *Canadian Journal of Forest Research-Revue Canadienne De Recherche Forestiere* **40** 1360-70

Cory R M, Ward C P, Crump B C and Kling G W 2014 Sunlight controls water column processing of carbon in arctic fresh waters *Science* **345** 925-8

de Groot W J, Cantin A S, Flannigan M D, Soja A J, Gowman L M and Newbery A 2013 A comparison of Canadian and Russian boreal forest fire regimes *Forest Ecology and Management* **294** 23-34

DeConto R M, Galeotti S, Pagani M, Tracy D, Schaefer K, Zhang T J, Pollard D and Beerling D J 2012 Past extreme warming events linked to massive carbon release from thawing permafrost (vol 484, pg 87, 2012) *Nature* **490** 292-

Epstein H E, Raynolds M K, Walker D A, Bhatt U S, Tucker C J and Pinzon J E 2012 Dynamics of aboveground phytomass of the circumpolar Arctic tundra during the past three decades *Environmental Research Letters* **7**

Falloon P D, Dankers R, Betts R A, Jones C D, Booth B B B and Lambert F H 2012 Role of vegetation change in future climate under the A1B scenario and a climate stabilisation scenario, using the HadCM3C Earth system model *Biogeosciences* **9** 4739-56

Flannigan M, Stocks B, Turetsky M and Wotton M 2009 Impacts of climate change on fire activity and fire management in the circumboreal forest *Glob. Change Biol.* **15** 549-60

French H M 1999 Past and present permafrost as an indicator of climate change *Polar Research* **18** 269-74

Gedalof Z e and Berg A A 2010 Tree ring evidence for limited direct CO2 fertilization of forests over the 20th century *Global Biogeochemical Cycles* **24** GB3027

Giglio L, Randerson J T, van der Werf G R, Kasibhatla P S, Collatz G J, Morton D C and DeFries R S 2010 Assessing variability and long-term trends in burned area by merging multiple satellite fire products *Biogeosciences* **7** 1171-86

Girardin M P, Hogg E H, Bernier P Y, Kurz W A, Guo X J and Cyr G 2015 Negative impacts of high temperatures on growth of black spruce forests intensify with the anticipated climate warming *Glob. Change Biol.* n/a-n/a

Guo L D, Ping C L and Macdonald R W 2007 Mobilization pathways of organic carbon from permafrost to arctic rivers in a changing climate *Geophysical Research Letters* **34**

Halpern B S, Walbridge S, Selkoe K A, Kappel C V, Micheli F, D'Agrosa C, Bruno J F, Casey K S, Ebert C, Fox H E, Fujita R, Heinemann D, Lenihan H S, Madin E M P, Perry M T, Selig E R, Spalding M, Steneck R and Watson R 2008 A global map of human impact on marine ecosystems *Science* **319** 948-52

Harden J W, Koven C D, Ping C-L, Hugelius G, David McGuire A, Camill P, Jorgenson T, Kuhry P, Michaelson G J, O'Donnell J A, Schuur E A G, Tarnocai C, Johnson K and Grosse G 2012 Field information links permafrost carbon to physical vulnerabilities of thawing *Geophysical Research Letters* **39** L15704

Hayes D J, McGuire A D, Kicklighter D W, Gurney K R, Burnside T J and Melillo J M 2011 Is the northern high-latitude land-based CO2 sink weakening? *Global Biogeochem. Cycles* **25** GB3018

Hewitt R, Bennett A, Breen A, Hollingsworth T, Taylor D L, Chapin F S, III and Rupp T S 2015 Getting to the root of the matter: landscape implications of plant-fungal interactions for tree migration in Alaska *Landscape Ecol* 1-17

Hickler T, Smith B, Prentice I C, Mjofors K, Miller P, Arneth A and Sykes M T 2008 CO(2) fertilization in temperate FACE experiments not representative of boreal and tropical forests *Glob. Change Biol.* **14** 1531-42

Holmes R M, McClelland J W, Peterson B J, Tank S E, Bulygina E, Eglinton T I, Gordeev V V, Gurtovaya T Y, Raymond P A, Repeta D J, Staples R, Striegl R G, Zhulidov A V and Zimov S A 2012 Seasonal and Annual Fluxes of Nutrients and Organic Matter from Large Rivers to the Arctic Ocean and Surrounding Seas *Estuaries and Coasts* **35** 369-82

Hugelius G, Strauss J, Zubrzycki S, Harden J W, Schuur E A G, Ping C L, Schirrmeister L, Grosse G, Michaelson G J, Koven C D, O'Donnell J A, Elberling B, Mishra U, Camill P, Yu Z, Palmtag J and Kuhry P 2014 Estimated stocks of circumpolar permafrost carbon with quantified uncertainty ranges and identified data gaps *Biogeosciences* **11** 6573-93

Hyvonen R, Agren G I, Linder S, Persson T, Cotrufo M F, Ekblad A, Freeman M, Grelle A, Janssens I A, Jarvis P G, Kellomaki S, Lindroth A, Loustau D, Lundmark T, Norby R J, Oren R, Pilegaard K, Ryan M G, Sigurdsson B D, Stromgren M, van Oijen M and Wallin G 2007 The likely impact of elevated CO2 , nitrogen deposition, increased temperature and management on carbon sequestration in temperate and boreal forest ecosystems: a literature review *New Phytol* **173** 463-80

Idso C D, Idso S B, Carter R M and Singer S F 2014 Climate Change Reconsidered: Biological Impacts. In: *Report of the Nongovernmental International Panel on Climate Change (NIPCC)* ed H Institute (Chicago, Illinois, U.S.A.: Heartland Institute)

Javeline D, Hellmann J J, Cornejo R C and Shufeldt G 2013 Expert Opinion on Climate Change and Threats to Biodiversity *Bioscience* **63** 666-73

Joly J L, Reynolds J and Robards M 2010 Recognizing When the "Best Scientific Data Available" Isn't *Stanford Environmental Law Journal* **29** 247-82

Jorgenson M T, Romanovsky V, Harden J, Shur Y, O'Donnell J, Schuur E A G, Kanevskiy M and Marchenko S 2010 Resilience and vulnerability of permafrost to climate change *Canadian Journal of Forest Research-Revue Canadienne De Recherche Forestiere* **40** 1219-36

Jorgenson T, Harden J, Kanevskiy M, O’Donnell J, Wickland K, Ewing S, Manies K, Zhuang Q, Shur Y, Striegl R and Koch J 2013 Reorganization of vegetation, hydrology and soil carbon after permafrost degradation across heterogeneous boreal landscapes *Environmental Research Letters* **8**

Kicklighter D W, Hayes D J, McClelland J W, Peterson B J, McGuire A D and Melillo J M 2013 Insights and issues with simulating terrestrial DOC loading of Arctic river networks *Ecological Applications* **23** 1817-36

Kloster S, Mahowald N M, Randerson J T and Lawrence P J 2012 The impacts of climate, land use, and demography on fires during the 21st century simulated by CLM-CN *Biogeosciences* **9** 509-25

Koven C D, Lawrence D M and Riley W J 2015a Permafrost carbon−climate feedback is sensitive to deep soil carbon decomposability but not deep soil nitrogen dynamics *Proceedings of the National Academy of Sciences* **112** 3752-7

Koven C D, Ringeval B, Friedlingstein P, Ciais P, Cadule P, Khvorostyanov D, Krinner G and Tarnocai C 2011 Permafrost carbon-climate feedbacks accelerate global warming *P Natl Acad Sci USA* **108** 14769-74

Koven C D, Schuur E A G, Schädel C, Bohn T J, Burke E J, Chen G, Chen X, Ciais P, Grosse G, Harden J W, Hayes D J, Hugelius G, Jafarov E E, Krinner G, Kuhry P, Lawrence D M, MacDougall A H, Marchenko S S, McGuire A D, Natali S M, Nicolsky D J, Olefeldt D, Peng S, Romanovsky V E, Schaefer K M, Strauss J, Treat C C and Turetsky M 2015b A simplified, data-constrained approach to estimate the permafrost carbon–climate feedback *Phil. Trans. R. Soc. A* **373** 20140423

Kurz W A, Stinson G and Rampley G 2008 Could increased boreal forest ecosystem productivity offset carbon losses from increased disturbances? *Philosophical Transactions of the Royal Society B: Biological Sciences* **363** 2259-68

Laudon H, Buttle J, Carey S K, McDonnell J, McGuire K, Seibert J, Shanley J, Soulsby C and Tetzlaff D 2012 Cross-regional prediction of long-term trajectory of stream water DOC response to climate change *Geophysical Research Letters* **39** L18404

Lawrence D M, Koven C D, Swenson S C, Riley W J and Slater A G 2015 Permafrost thaw and resulting soil moisture changes regulate projected high-latitude CO 2 and CH 4 emissions *Environmental Research Letters* **10** 094011

Lawrence D M, Slater A G and Swenson S C 2012 Simulation of Present-Day and Future Permafrost and Seasonally Frozen Ground Conditions in CCSM4 *Journal of Climate* **25** 2207-25

Lenton T M, Held H, Kriegler E, Hall J W, Lucht W, Rahmstorf S and Schellnhuber H J 2008 Tipping elements in the Earth's climate system *P Natl Acad Sci USA* **105** 1786-93

Maaroufi N I, Nordin A, Hasselquist N J, Bach L H, Palmqvist K and Gundale M J 2015 Anthropogenic nitrogen deposition enhances carbon sequestration in boreal soils *Glob. Change Biol.* **21** 3169-80

MacDougall A H, Avis C A and Weaver A J 2012 Significant contribution to climate warming from the permafrost carbon feedback *Nature Geosci* **5** 719-21

Mack M C, Bret-Harte M S, Hollingsworth T N, Jandt R R, Schuur E A G, Shaver G R and Verbyla D L 2011 Carbon loss from an unprecedented Arctic tundra wildfire *Nature* **475** 489-92

McGuire A D, Anderson L G, Christensen T R, Dallimore S, Guo L, Hayes D J, Heimann M, Lorenson T D, Macdonald R W and Roulet N 2009 Sensitivity of the carbon cycle in the Arctic to climate change *Ecological Monographs* **79** 523-55

McGuire A D, Hayes D J, Kicklighter D W, Manizza M, Zhuang Q, Chen M, Follows M J, Gurney K R, McClelland J W, Melillo J M, Peterson B J and Prinn R G 2010 An analysis of the carbon balance of the Arctic Basin from 1997 to 2006 *Tellus B* **62** 455-74

Morgan M G 2014 Use (and abuse) of expert elicitation in support of decision making for public policy *Proceedings of the National Academy of Sciences* **111** 7176-84

Moss R H, Edmonds J A, Hibbard K A, Manning M R, Rose S K, van Vuuren D P, Carter T R, Emori S, Kainuma M, Kram T, Meehl G A, Mitchell J F B, Nakicenovic N, Riahi K, Smith S J, Stouffer R J, Thomson A M, Weyant J P and Wilbanks T J 2010 The next generation of scenarios for climate change research and assessment *Nature* **463** 747-56

Myers-Smith I H, Harden J W, Wilmking M, Fuller C C, McGuire A D and Chapin F S 2008 Wetland succession in a permafrost collapse: interactions between fire and thermokarst *Biogeosciences* **5** 1273-86

O'Donnell J A, Harden J W, McGuire A D, Kanevskiy M Z, Jorgenson M T and Xu X M 2011 The effect of fire and permafrost interactions on soil carbon accumulation in an upland black spruce ecosystem of interior Alaska: implications for post-thaw carbon loss *Glob. Change Biol.* **17** 1461-74

Pan Y, Birdsey R A, Fang J, Houghton R, Kauppi P E, Kurz W A, Phillips O L, Shvidenko A, Lewis S L, Canadell J G, Ciais P, Jackson R B, Pacala S, McGuire A D, Piao S, Rautiainen A, Sitch S and Hayes D 2011 A Large and Persistent Carbon Sink in the World's Forests *Science*

Payette S, Delwaide A, Caccianiga M and Beauchemin M 2004 Accelerated thawing of subarctic peatland permafrost over the last 50 years *Geophysical Research Letters* **31** L18208

Peñuelas J, Canadell J G and Ogaya R 2011 Increased water-use efficiency during the 20th century did not translate into enhanced tree growth *Global Ecology and Biogeography* **20** 597-608

Potter C S and Klooster S A 1997 Global model estimates of carbon and nitrogen storage in litter and soil pools: Response to changes in vegetation quality and biomass allocation *Tellus B* **49** 1-17

Qian H, Joseph R and Zeng N 2010 Enhanced terrestrial carbon uptake in the Northern High Latitudes in the 21st century from the Coupled Carbon Cycle Climate Model Intercomparison Project model projections *Glob. Change Biol.* **16** 641-56

Raynolds M K, Walker D A, Epstein H E, Pinzon J E and Tucker C J 2012 A new estimate of tundra-biome phytomass from trans-Arctic field data and AVHRR NDVI *Remote Sensing Letters* **3** 403-11

Rocha A V, Loranty M M, Higuera P E, Mack M C, Hu F S, Jones B M, Breen A L, Rastetter E B, Goetz S J and Shaver G R 2012 The footprint of Alaskan tundra fires during the past half-century: implications for surface properties and radiative forcing *Environmental Research Letters* **7**

Rupp T S, Starfield A M and Chapin Iii F S 2000 A frame-based spatially explicit model of subarctic vegetation response to climatic change: comparison with a point model *Landscape Ecol* **15** 383-400

Saugier B, Roy J and Mooney H A 2001 *Terrestrial Global Productivity,* ed J Roy*, et al.* (San Diego: Academic Press) pp 543-57

Schaefer K, Lantuit H, Romanovsky V E, Schuur E A G and Witt R 2014 The impact of the permafrost carbon feedback on global climate *Environmental Research Letters* **9** 085003

Schaefer K, Zhang T, Bruhwiler L and Barrett A P 2011 Amount and timing of permafrost carbon release in response to climate warming *Tellus B* no-no

Schirrmeister L, Siegert C, Kuznetsova T, Kuzmina S, Andreev A, Kienast F, Meyer H and Bobrov A 2002 Paleoenvironmental and paleoclimatic records from permafrost deposits in the Arctic region of Northern Siberia *Quatern Int* **89** 97-118

Schuur E A G, Abbott B W, Bowden W B, Brovkin V, Camill P, Canadell J G, Chanton J P, Chapin F S, III, Christensen T R, Ciais P, Crosby B T, Czimczik C I, Grosse G, Harden J, Hayes D J, Hugelius G, Jastrow J D, Jones J B, Kleinen T, Koven C D, Krinner G, Kuhry P, Lawrence D M, McGuire A D, Natali S M, O’Donnell J A, Ping C L, Riley W J, Rinke A, Romanovsky V E, Sannel A B K, Schädel C, Schaefer K, Sky J, Subin Z M, Tarnocai C, Turetsky M R, Waldrop M P, Walter Anthony K M, Wickland K P, Wilson C J and Zimov S A 2013 Expert assessment of vulnerability of permafrost carbon to climate change *Climatic Change* 1-16

Schuur E A G, McGuire A D, Schadel C, Grosse G, Harden J W, Hayes D J, Hugelius G, Koven C D, Kuhry P, Lawrence D M, Natali S M, Olefeldt D, Romanovsky V E, Schaefer K, Turetsky M R, Treat C C and Vonk J E 2015 Climate change and the permafrost carbon feedback *Nature* **520** 171-9

Shur Y L and Jorgenson M T 2007 Patterns of permafrost formation and degradation in relation to climate and ecosystems *Permafrost and Periglacial Processes* **18** 7-19

Sturm M, McFadden J P, Liston G E, Chapin F S, Racine C H and Holmgren J 2001 Snow-shrub interactions in Arctic tundra: A hypothesis with climatic implications *Journal of Climate* **14** 336-44

Sutherland W J, Freckleton R P, Godfray H C J, Beissinger S R, Benton T, Cameron D D, Carmel Y, Coomes D A, Coulson T, Emmerson M C, Hails R S, Hays G C, Hodgson D J, Hutchings M J, Johnson D, Jones J P G, Keeling M J, Kokko H, Kunin W E, Lambin X, Lewis O T, Malhi Y, Mieszkowska N, Milner-Gulland E J, Norris K, Phillimore A B, Purves D W, Reid J M, Reuman D C, Thompson K, Travis J M J, Turnbull L A, Wardle D A and Wiegand T 2013 Identification of 100 fundamental ecological questions *J Ecol* **101** 58-67

Thornton P E, Lamarque J-F, Rosenbloom N A and Mahowald N M 2007 Influence of carbon-nitrogen cycle coupling on land model response to CO2 fertilization and climate variability *Global Biogeochemical Cycles* **21** GB4018

Treat C C and Frolking S 2013 Carbon Storage: A permafrost carbon bomb? *Nature Clim. Change* **3** 865-7

UNEP 2013 The Emissions Gap Report 2013. United Nations Environment Programme (UNEP). (Nairobi

van der Werf G R, Randerson J T, Giglio L, Collatz G J, Mu M, Kasibhatla P S, Morton D C, DeFries R S, Jin Y and van Leeuwen T T 2010 Global fire emissions and the contribution of deforestation, savanna, forest, agricultural, and peat fires (1997–2009) *Atmos. Chem. Phys.* **10** 11707-35

Vonk J E, Sanchez-Garcia L, van Dongen B E, Alling V, Kosmach D, Charkin A, Semiletov I P, Dudarev O V, Shakhova N, Roos P, Eglinton T I, Andersson A and Gustafsson O 2012 Activation of old carbon by erosion of coastal and subsea permafrost in Arctic Siberia *Nature* **489** 137-40

Whiteman G, Hope C and Wadhams P 2013 Climate science: Vast costs of Arctic change *Nature* **499** 401-3

Wieczorek S, Ashwin P, Luke C M and Cox P M 2011 Excitability in ramped systems: the compost-bomb instability *P Roy Soc a-Math Phy* **467** 1243-69

Yarie J and Van Cleve K 2010 Long-term monitoring of climatic and nutritional affects on tree growth in interior Alaska *Canadian Journal of Forest Research-Revue Canadienne De Recherche Forestiere* **40** 1325-35

Zachos J C, Dickens G R and Zeebe R E 2008 An early Cenozoic perspective on greenhouse warming and carbon-cycle dynamics *Nature* **451** 279-83

Zhang X, He J, Zhang J, Polyakov I, Gerdes R, Inoue J and Wu P 2013 Enhanced poleward moisture transport and amplified northern high-latitude wetting trend *Nature Clim. Change* **3** 47-51

Zickfeld K, Morgan M G, Frame D J and Keith D W 2010 Expert judgments about transient climate response to alternative future trajectories of radiative forcing *P Natl Acad Sci USA* **107** 12451-6